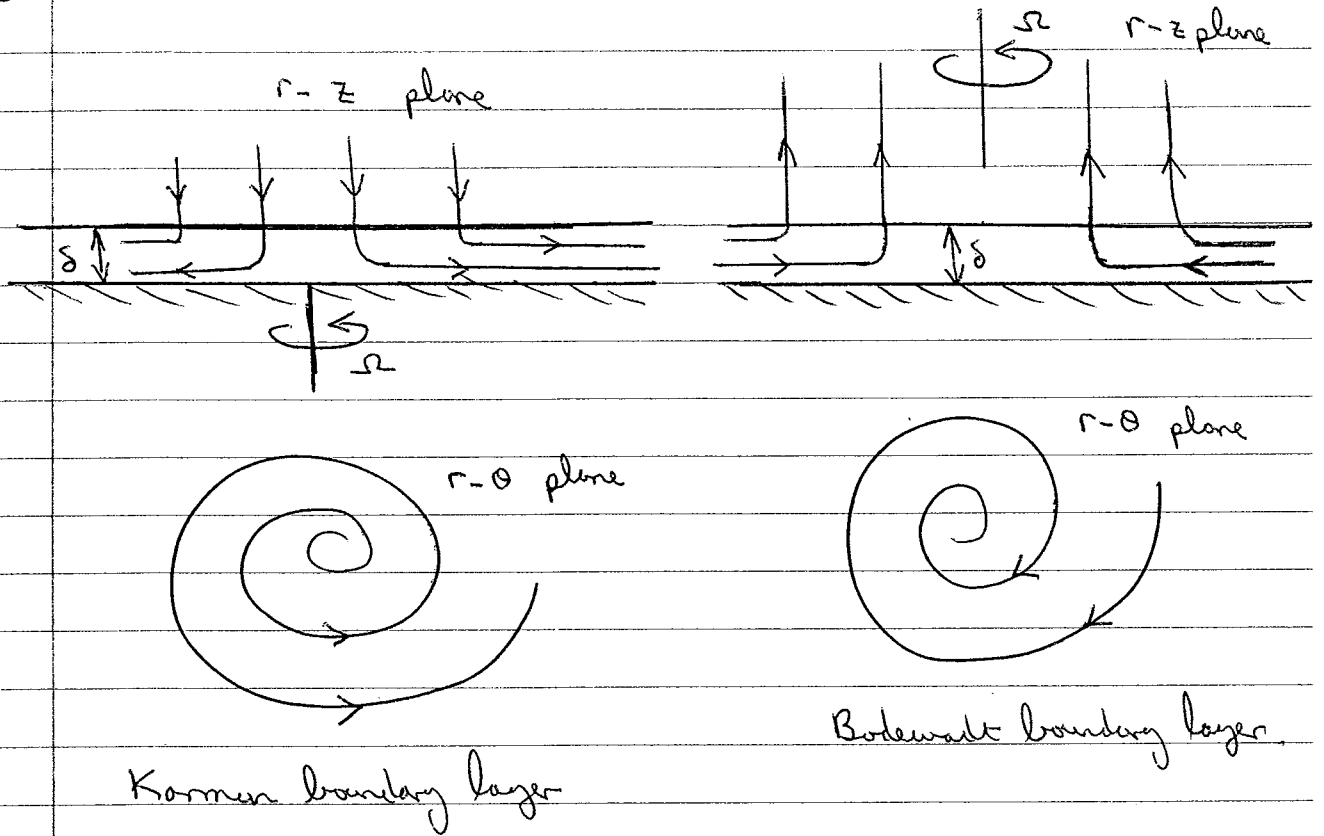


① (a)



Outside the boundary layer there is a balance between

$\frac{\partial p}{\partial r}$ and $\rho u^2/r$. This sets up a radial pressure gradient

which is imposed on the boundary. In the case of a Bodewadt

layer there is a low pressure at $r=0$ and high pressure at large

r . For a Kármán layer $\frac{\partial p}{\partial r} = 0$. In the boundary

layer this balance is disturbed because u_0 varies with z . In

the Kármán layer the fluid is centrifuged outward and in the

Bodewadt layer u_0 is diminished so $\frac{\partial p}{\partial r}$ pushes the fluid inward. [30%]

(b) In Kármán layer $\partial p / \partial r = 0$ and the balance is between $\rho u_0^2 / r$ and the viscous stress $\rho \nu \partial^2 u_r / \partial z^2$.

Thus,

$$\rho \frac{u_0^2}{r} \sim \rho \nu \frac{u_r}{\delta^2}$$

In the boundary layer $\partial p / \partial r \sim \rho u_0^2 / r$ in core,

and the pressure gradient $\partial p / \partial r$ in the boundary layer

is balanced by the viscous stress $\rho \nu \partial^2 u_r / \partial z^2$:

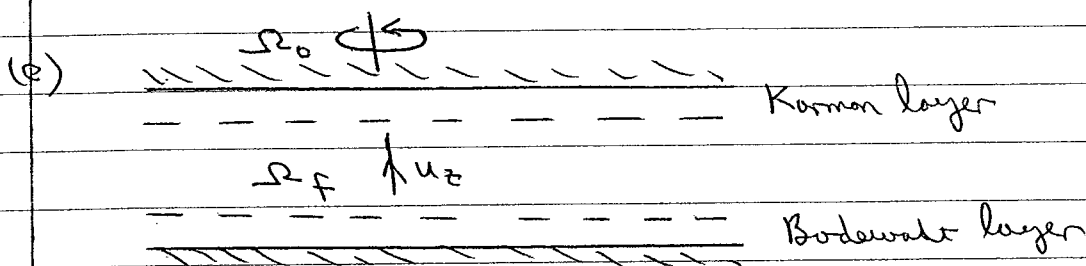
$$\frac{\partial p}{\partial r} \sim \left(\rho \frac{u_0^2}{r} \right)_{\text{core}} \sim \rho \nu \frac{u_r}{\delta^2}$$

In both cases $u_0 \sim \Omega r$ and $u_0 \sim u_r$, so,

$$\rho \frac{(\Omega r)^2}{r} \sim \rho \nu \frac{\Omega r}{\delta^2}$$

$$\Rightarrow \delta \sim \sqrt{\nu / \Omega}$$

[35%]



$$u_z = 1.4 \sqrt{\nu \Omega R_f} = 0.9 \sqrt{\nu (\Omega_0 - \Omega_f)}$$

$$\Rightarrow \left(\frac{1.4}{0.9} \right)^2 \Omega_f = \Omega_0 - \Omega_f$$

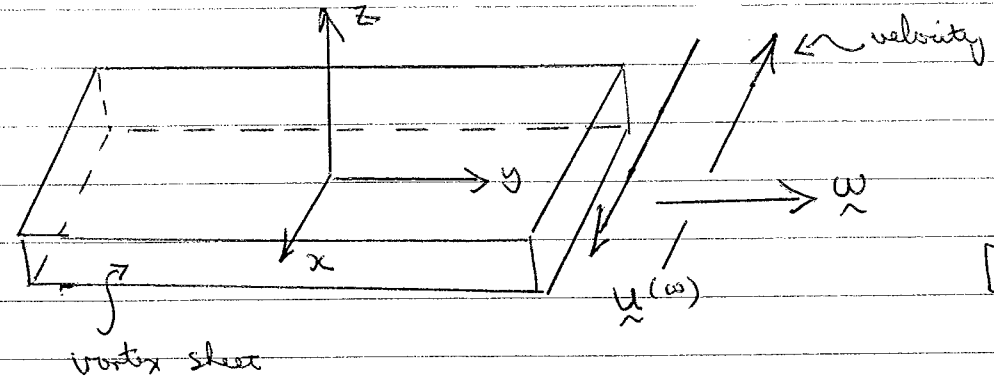
$$\Rightarrow \underline{\underline{\Omega_f = 0.292 \Omega_0}}$$

[35%]

Examiner's comment:

The answers were very good by and large, with the occasional hopeless attempt. Most marks were lost in the order-of-magnitude part of the question. On reflection the question needed a sting in the tail.

(3) (a)



[10%]

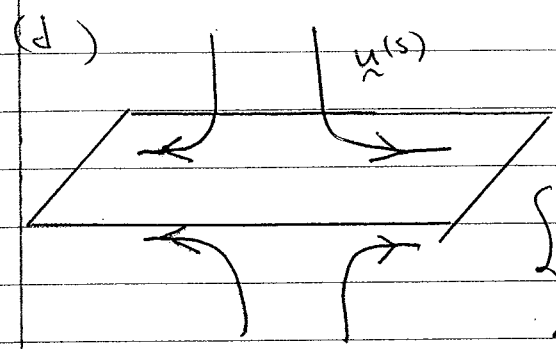
(b)
$$\vec{\omega} = \begin{vmatrix} \hat{i} & \hat{j} & \hat{k} \\ 0 & 0 & \frac{\partial}{\partial z} \\ u_x & 0 & 0 \end{vmatrix} = \frac{\partial u_x}{\partial z} \hat{j} \Rightarrow \omega_y = \frac{\partial u_x}{\partial z}$$

$$\left\{ \begin{aligned} \text{Flux} &= \int_{-\infty}^{\infty} \omega_y dz = \int_{-\infty}^{\infty} \frac{\partial u_x}{\partial z} dz = \Delta u_x \text{ (velocity jump)} \\ \text{Flux} &= \int_{-\infty}^{\infty} \omega_y dz = \frac{\Phi}{\sqrt{\pi} \delta} \int_{-\infty}^{\infty} e^{-z^2/\delta^2} dz = \frac{\Phi}{\sqrt{\pi}} \int_{-\infty}^{\infty} e^{-s^2} ds = \Phi \end{aligned} \right.$$

[20%]

(c)
$$(\vec{\omega} \cdot \nabla) u_x^{(ω)} = \omega_y \frac{\partial}{\partial y} u_x^{(ω)} = 0$$

$$(u_x^{(ω)} \cdot \nabla) \vec{\omega} = u_x^{(ω)} \frac{\partial}{\partial x} \vec{\omega} = 0 \quad [10%]$$



$$u^{(s)} = (0, \alpha y, -\alpha z)$$

$$\left\{ \begin{aligned} u^{(s)} \cdot \nabla \vec{\omega} &= -\alpha z \frac{\partial}{\partial z} \omega_y \hat{j} \\ \vec{\omega} \cdot \nabla u^{(s)} &= \omega_y \frac{\partial}{\partial y} (\alpha y \hat{j}) \end{aligned} \right.$$

$$u^{(s)} \cdot \nabla \vec{\omega} = \vec{\omega} \cdot \nabla u^{(s)} + \nu \nabla^2 \vec{\omega}$$

$$\Rightarrow -\alpha z \frac{\partial \omega_y}{\partial z} = \omega_y \alpha + \nu \frac{\partial^2 \omega_y}{\partial z^2}$$

$$\Rightarrow -\alpha z \left[-\frac{2z}{\delta^2} \right] \omega_y = \alpha \omega_y + \nu \frac{\partial}{\partial z} \left[-\frac{2z}{\delta^2} \omega_y \right]$$

$$= \alpha \omega_y + \nu \left[-\frac{2}{\delta^2} \omega_y + \frac{4z^2}{\delta^4} \omega_y \right]$$

$$\Rightarrow \alpha \omega_y \left[\frac{2z^2}{\delta^2} - 1 \right] = \frac{2\nu}{\delta^2} \omega_y \left[\frac{2z^2}{\delta^2} - 1 \right]$$

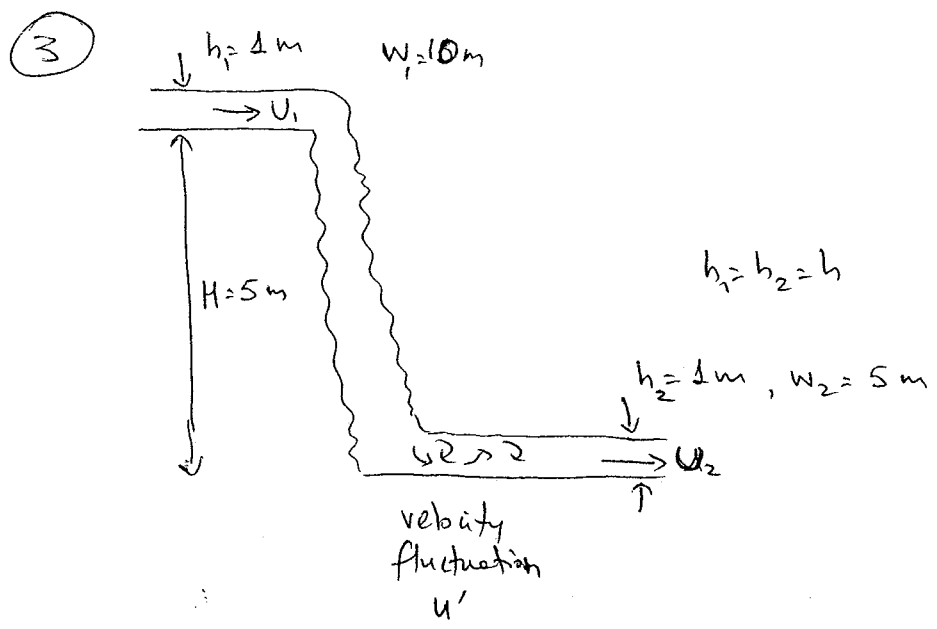
$$\Rightarrow \alpha = \frac{2\nu}{\delta^2}$$

$$\Rightarrow \underline{\underline{\delta = \sqrt{2\nu/\alpha}}} \quad [40\%]$$

(e) The inward advection of vorticity counters the tendency for vorticity to diffuse outward. The stretching of vorticity maintains the vorticity at constant strength despite diffusion [20%]

Examiner's comment:

The answers were mostly sound, though curiously most marks were lost in part (b), which should have been easy.



(a) Continuity: $U_1 h W_1 = U_2 h W_2 \Rightarrow U_2 = 2 \text{ m/s}$

Total energy: $\frac{1}{2} \rho U_1^2 + \rho g H = \frac{1}{2} \rho U_2^2 + \left(\frac{3}{2}\right) \rho u'^2$ IMPORTANT!!

$\Rightarrow 3u'^2 = 2gH - (U_2^2 - U_1^2) \Rightarrow u' = 5.63 \text{ m/s}$

Turbulence "picks up" the energy from the fall (potential energy) that the mean flow didn't pick up. Note that h_2 to be equal to the one before the waterfall is probably an oversimplification of reality. (30%)

(b) A reasonable integral lengthscale is h (or a fraction of h). W_2 is not a good estimate, since it is the smallest dimension of the flow that usually serves as the order of magnitude of the turbulent lengthscale.

So: $Re_t = \frac{u' h}{\nu} = 5.6 \times 10^6$ (2) $\nu_{\text{water}} = 10^{-6} \text{ m}^2/\text{s}$ at 20°C

$\epsilon = u'^3 / h = 178.5 \text{ m}^2/\text{s}^3$

$\lambda_k = h \cdot Re_t^{-3/4} = 9 \mu\text{m}$

(40%)

(c) At the foot of the waterfall, $u' \sim 5.6$ m/s and the mean velocity is $U_2 = 2$ m/s.

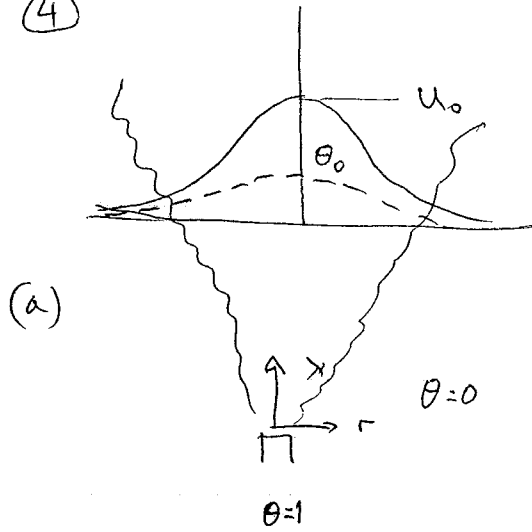
Taylor's assumption states that $\frac{\partial}{\partial t} = U \frac{\partial}{\partial x}$

or that the turbulence is "frozen" as it is convected past an observer. It is valid if $u' \ll U_2$ which is not satisfied in our case.

Upstream there is no turbulence so Taylor's hypothesis is ok (albeit trivial). Downstream, the turbulence will eventually decay and will become $\ll U_2$. If the flow in the river becomes something like a flat plate boundary layer, Taylor's hypothesis will be reasonable far from the wall, not very good close to the wall.

(30%)

4



$$U(x,r) = U_0(x) \cdot G(\eta)$$

$$\eta = r/\delta$$

$\delta =$ unique function of x
 $=$ characteristic width of jet

All Reynolds stresses, fluxes etc
 can be written as
 $F_0(x) \cdot F(\eta)$

For the mean scalar, $\theta(x,r) = \theta_0(x) \cdot G_\theta(\eta)$

For the turbulence, $\overline{u^2}(x,r) = u_0^2(x) \cdot G_{uu}(\eta)$

For the scalar variance, $\overline{\theta^2}(x,r) = \theta_0^2(x) \cdot G_{\theta\theta}(\eta)$

Self-preservation implies also :

- (i) $\theta_0^2/\theta_0^2, u_0^2/U_0^2$ to be constant (independent of x)
- (ii) turbulence structure does not change (e.g. balance of terms in κ or Re-stress transport equation stays the same)

§

(50%)

(b) For the round jet $\delta \sim x$.

From conservation of momentum, $M = \text{const}$

$$M = \int_0^\infty 2\pi r U^2(r) dr = \int_0^\infty 2\pi \eta \delta U_0^2(x) G^2(\eta) \delta d\eta$$

$$= \delta^2 U_0^2 \underbrace{\int_0^\infty 2\pi \eta G^2(\eta) d\eta}_{\text{indep. of } x} = \text{const}$$

$$\Rightarrow \delta^2 U_0^2 = \text{ind. of } x \Rightarrow U_0 \sim x^{-1}$$

$$\Rightarrow u_0 \sim x^{-1} \text{ since } u_0/U_0 = \text{const}$$

From conservation of scalar :

$$\int_0^\infty 2\pi r U(r) \theta(r) dr = \text{const}$$

$$\Rightarrow \int_0^\infty 2\pi r_0 \delta U_0(x) G(x) \theta_0(x) G_0(x) \delta dx = \text{const}$$

$\Rightarrow \delta^2 U_0 \theta_0$ must be independent of x

$$\Rightarrow \theta_0 \sim x^{-1}$$

$$\Rightarrow \delta \sim x^{-1}$$

Scalar dissipation is modelled as

$$N = C \cdot \frac{g^2}{T_{turb}}, \quad T_{turb} = \frac{L_{turb}}{u'}$$

$$= \frac{\delta}{u_0} \text{ for this problem}$$

$$\Rightarrow N \sim \frac{g_0^2 u_0}{\delta} \sim x^{-4}$$

Scalar dissipation decays very quickly downstream.

(50%)